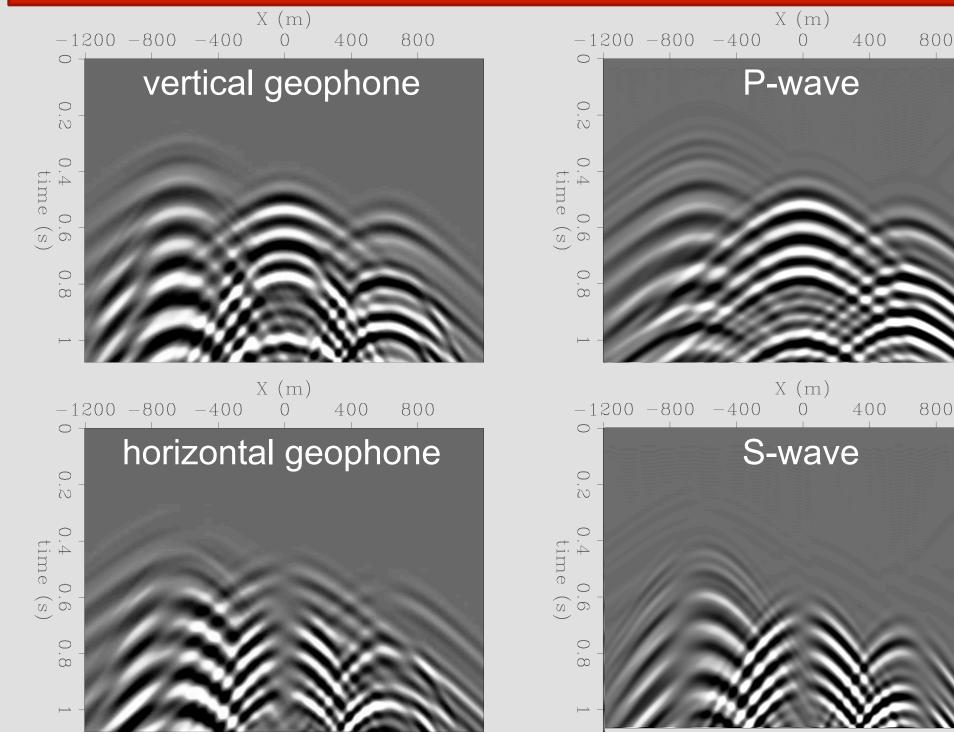
P/S separation of ocean-bottom seismic data by inversion in a homogeneous medium

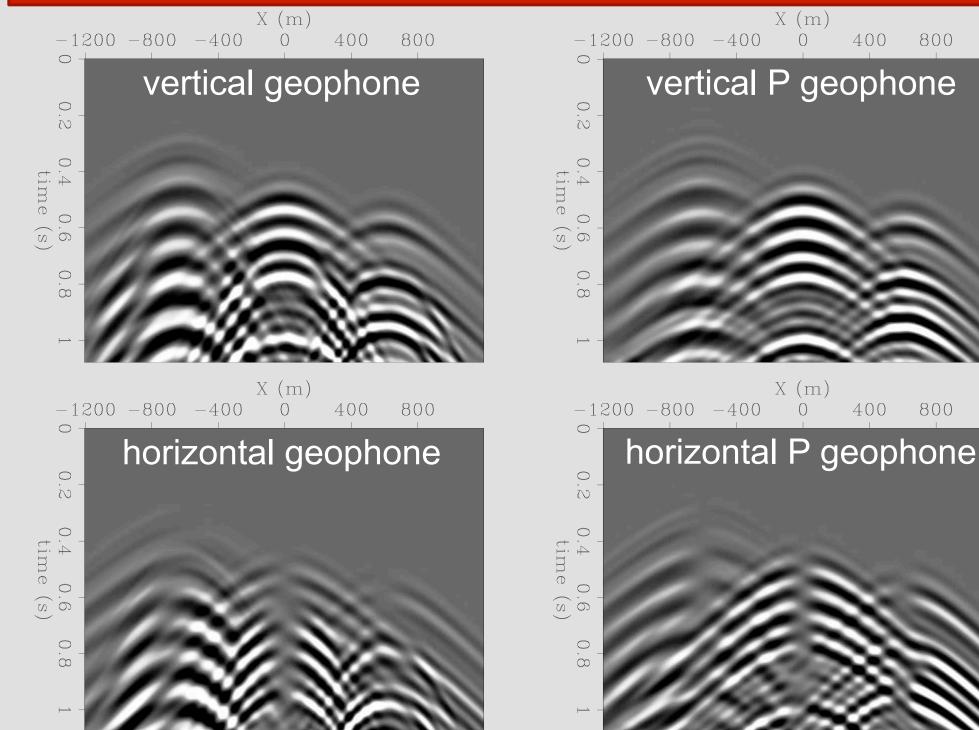
SEP sponsor's meeting May 23, 2012

Ohad Barak SEP 147, p. 261

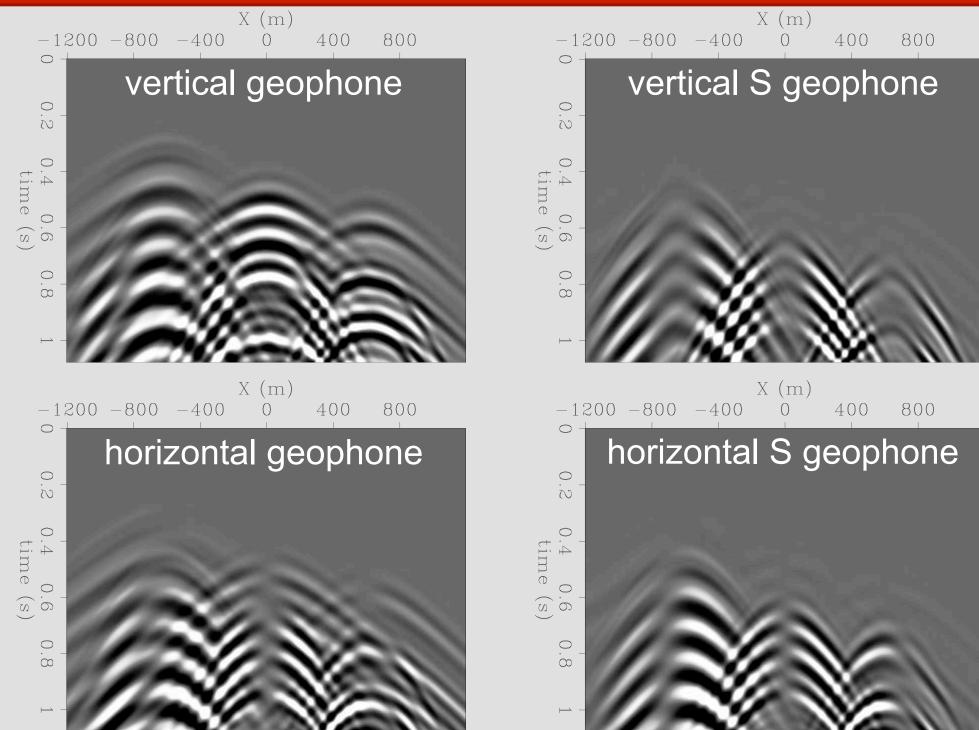
P/S separation



P/S separation



P/S separation



<u>Content</u>

- 1. Motivation
- 2. Theory
 - General concept
 - Inversion setup
- 3. Synthetic examples
 - Land
 - Ocean-bottom seismic
- 4. Conclusions
- 5. Road ahead

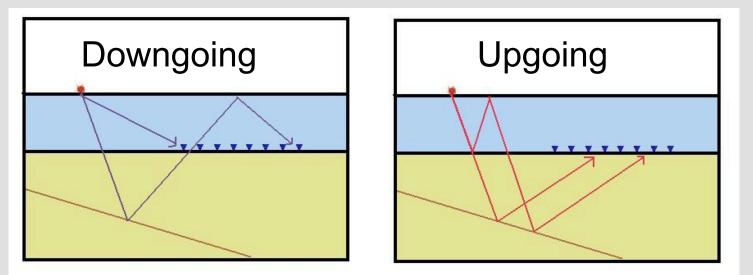
<u>Content</u>

1. Motivation

2. Theory

- General concept
- Inversion setup
- 3. Synthetic examples
 - Land
 - Ocean-bottom seismic
- 4. Conclusions
- 5. Road ahead

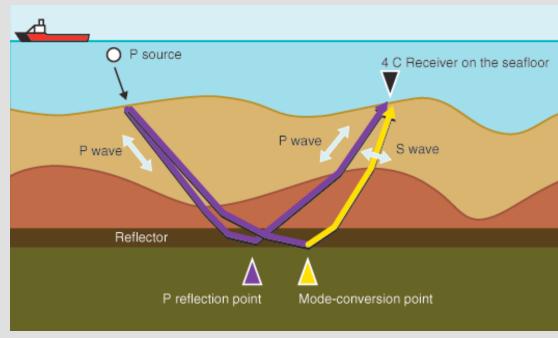
P/S separation of ocean-bottom seismic data



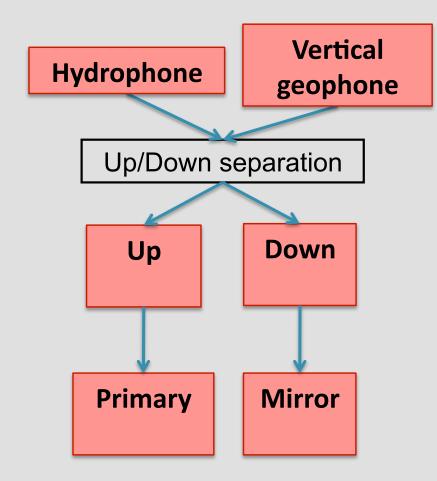
Wong and Ronen, SEP-138

Four-component acquisition

- 1 Hydrophone: pressure
- 3 Geophones: vertical and horizontal particle motion

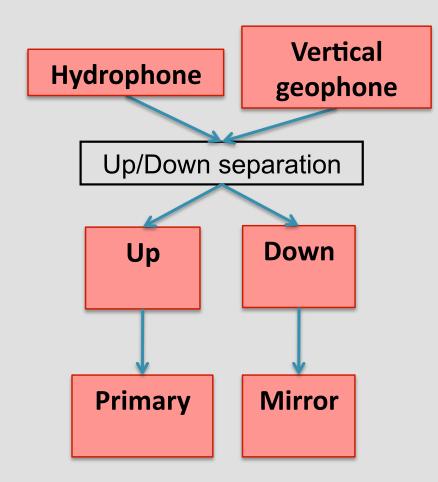


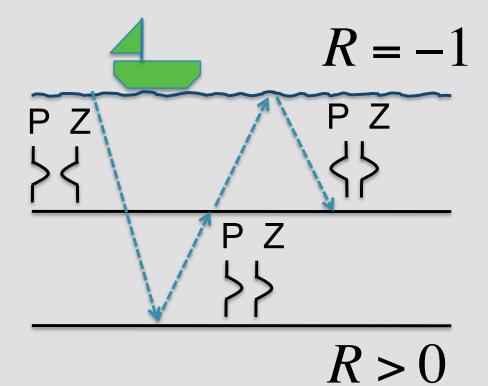
1. Assume vertical geophone is P.



1. Assume vertical geophone is P.

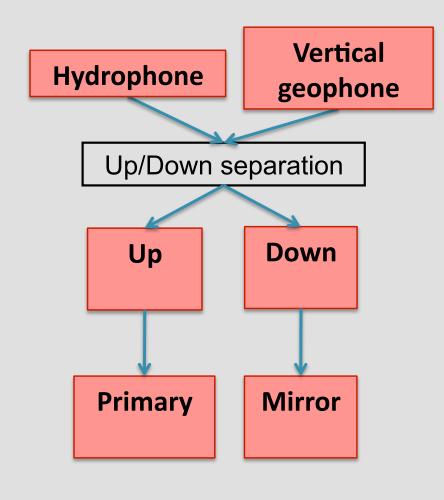
P-Z summation

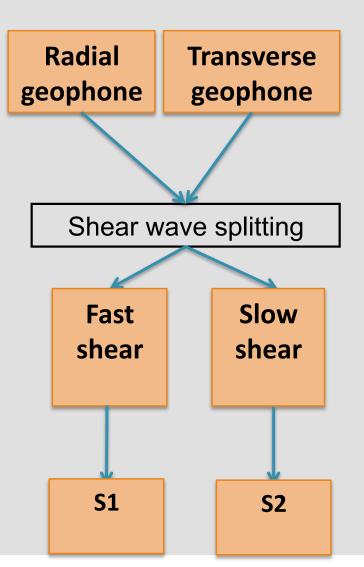




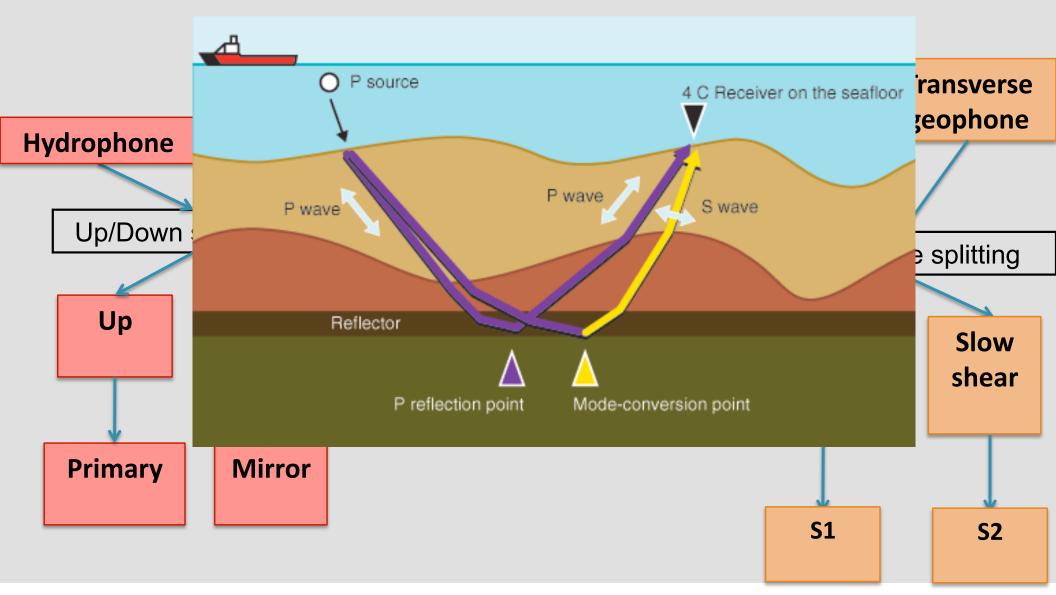
1. Assume vertical geophone is P.

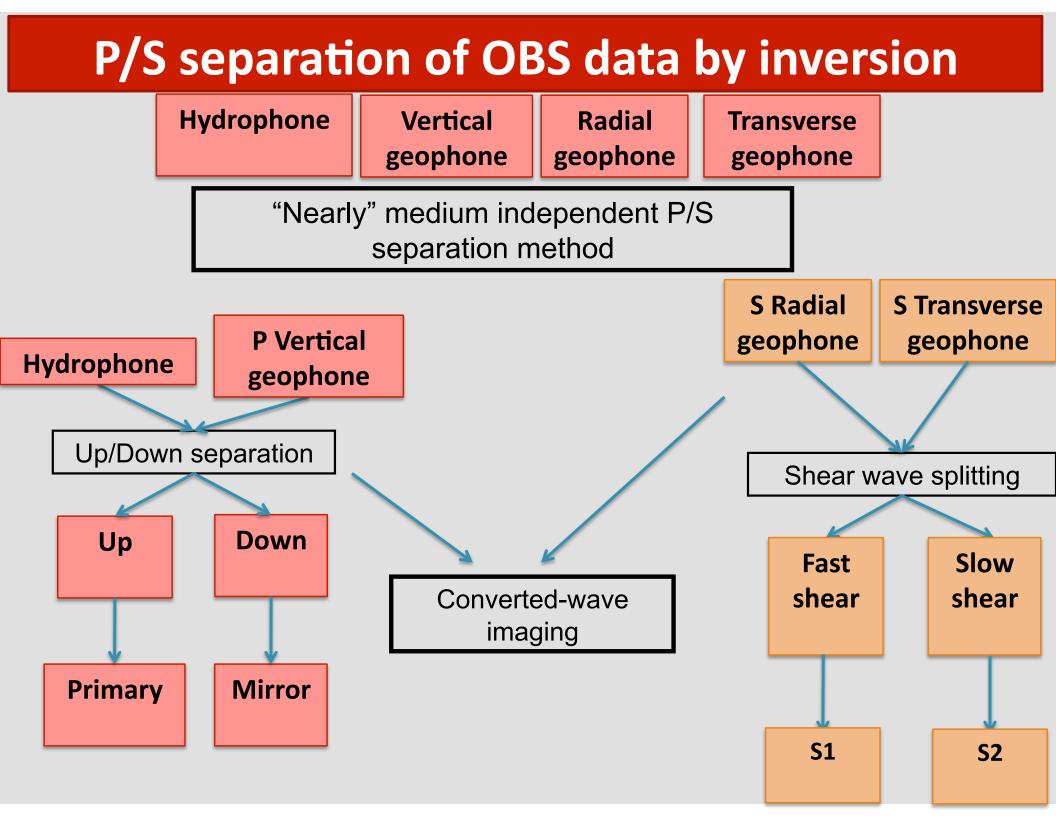
2. Assume horizontal geophones are S.

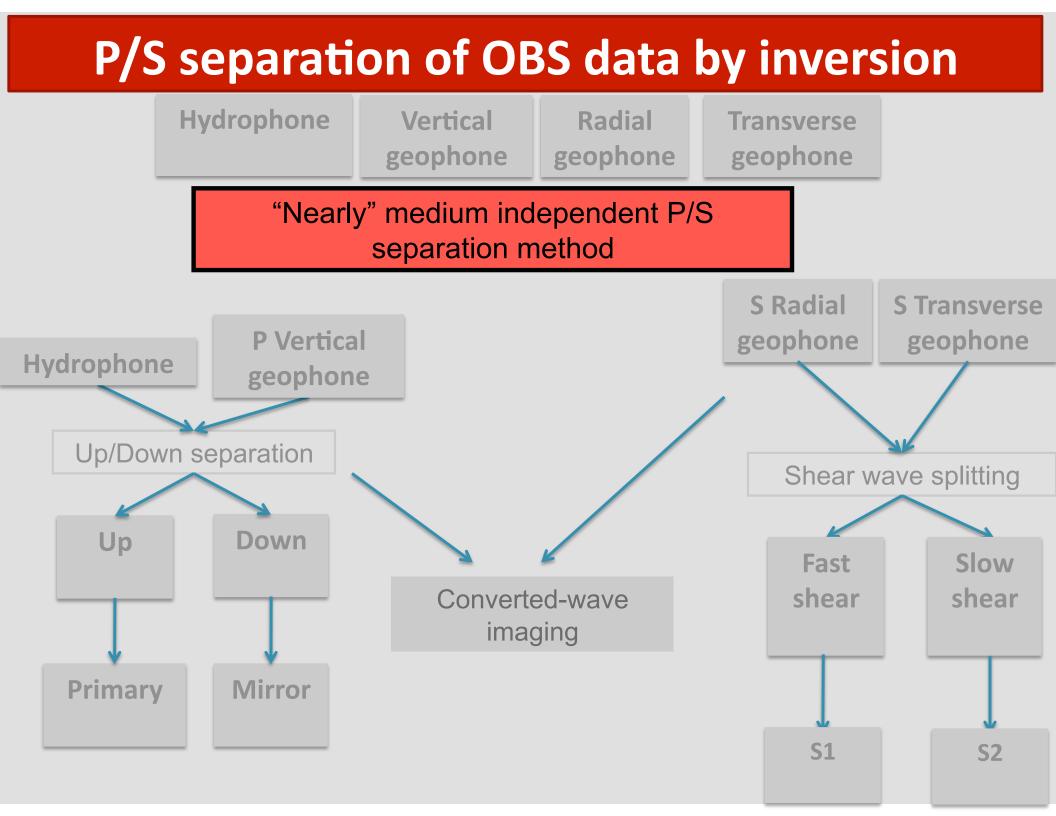




 Assume vertical geophone is P. 2. Assume horizontal geophones are S.



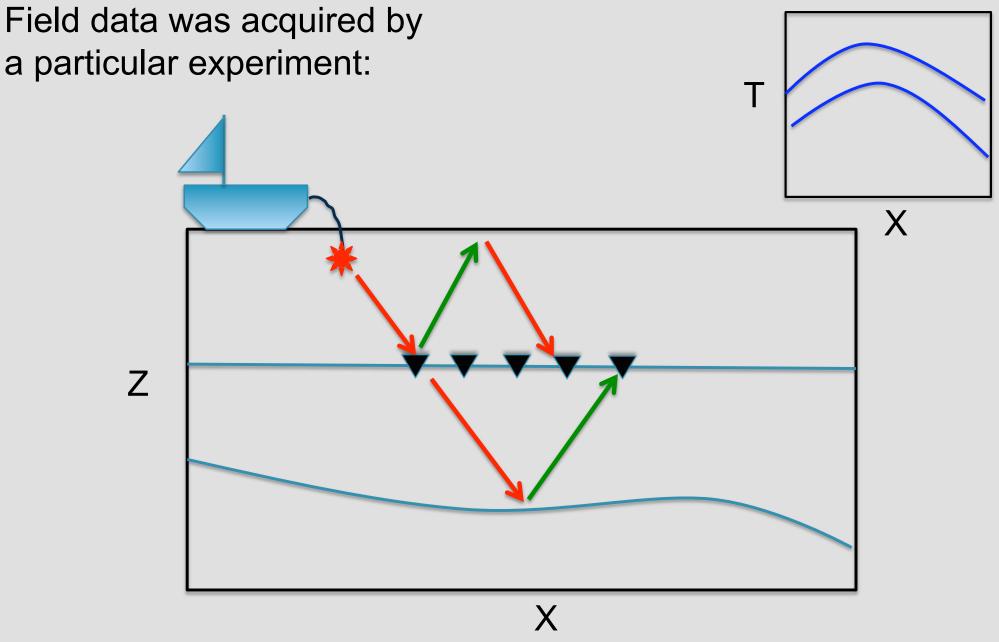




<u>Content</u>

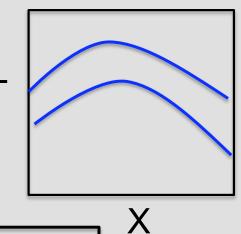
1. Motivation

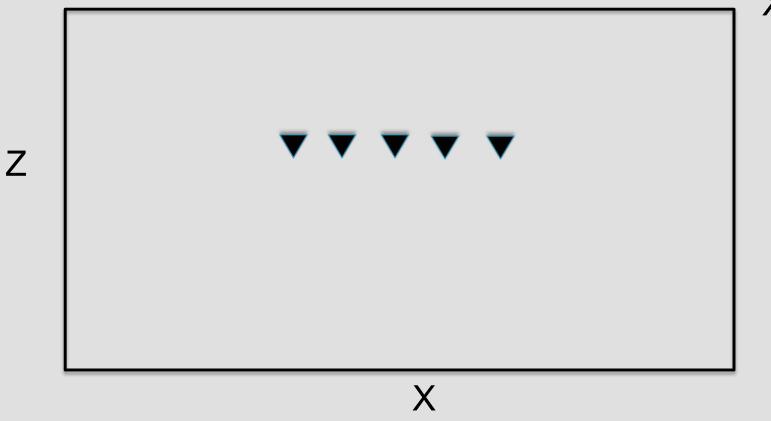
- 2. Theory
 - General concept.
 - Inversion setup.
- 3. Synthetic examples
 - Land.
 - Ocean-bottom seismic.
- 4. Conclusions.
- 5. Road ahead.



The same field data could have been the result of a different experiment.

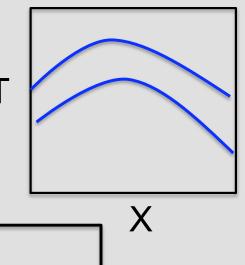
- 1. Different medium
- 2. Different acquisition

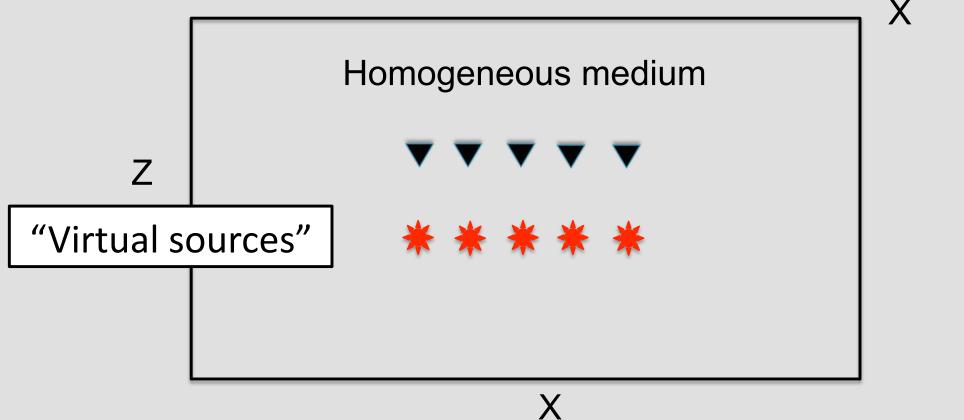




The same field data could have been the result of a different experiment.

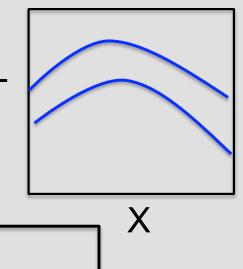
- 1. Different medium
- 2. Different acquisition

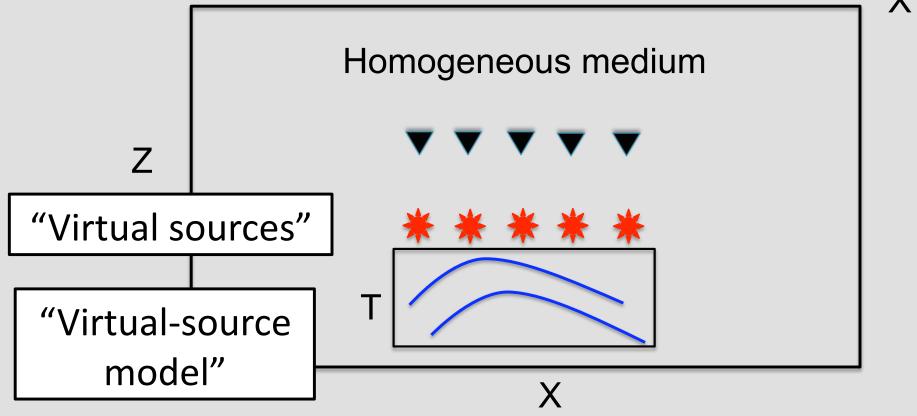




The same field data could have been the result of a different experiment.

- 1. Different medium
- 2. Different acquisition



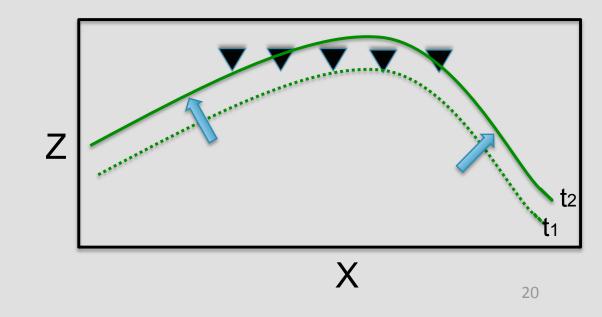


$$d_{obs}(x_r) = w(\omega)G(\omega, x_s, x_r)$$

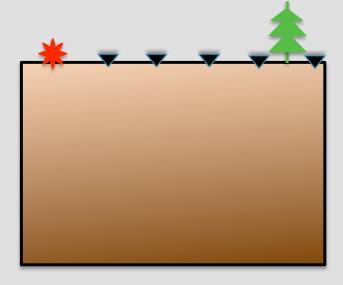
$$\widehat{\bigcup}$$

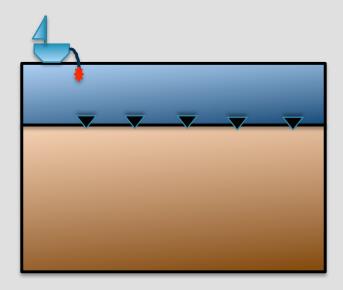
$$d_{equiv}(x_r) = \sum_{x'_s} f(\omega, x'_s) \tilde{G}(\omega, x'_s, x_r)$$

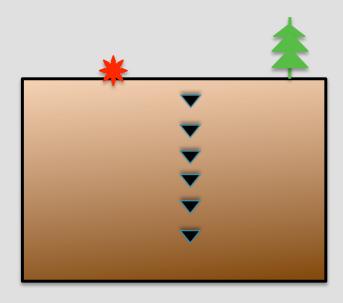
• Receiver data = a sampling of a field over time.

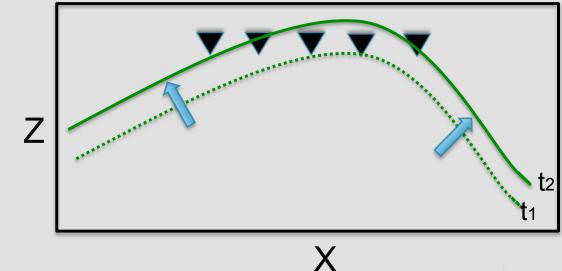


• Receiver data = a sampling of a field over time.

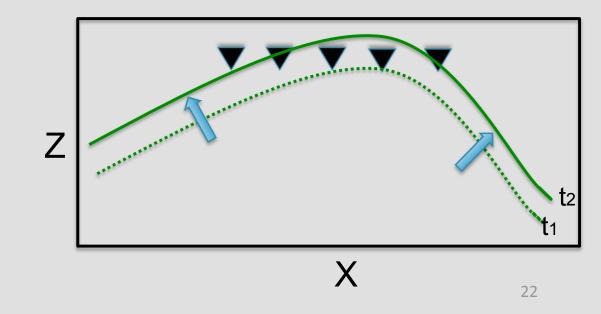






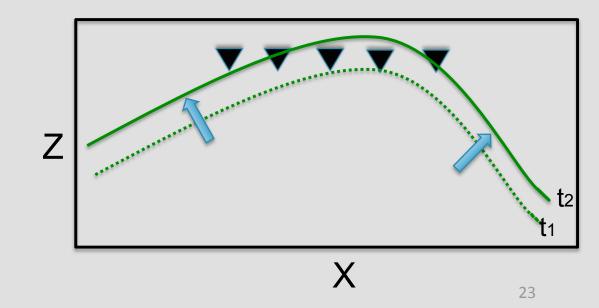


- Receiver data = a sampling of a field over time.
- Matching the recorded data = matching the wavefield at the receiver locations.



- Receiver data = a sampling of a field over time.
- Matching the recorded data = matching the wavefield at the receiver locations.
- Assuming wrong medium parameters:

How accurate is the reconstructed wavefield value one depth level <u>BELOW</u> the receivers?

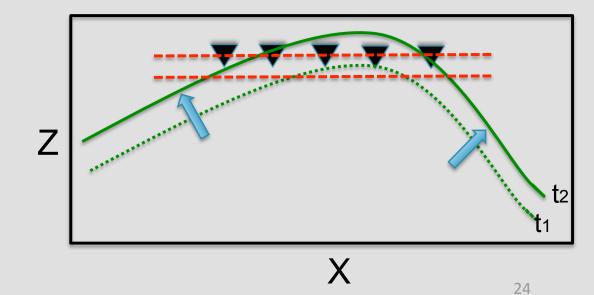


- Receiver data = a sampling of a field over time.
- Matching the recorded data = matching the wavefield at the receiver locations.
- Assuming wrong medium parameters:

How accurate is the reconstructed wavefield value one depth level <u>BELOW</u> the receivers?

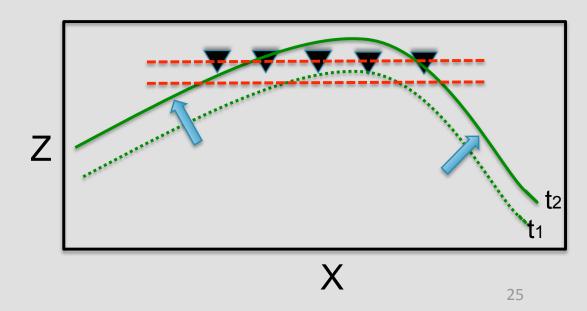
- 1. Typical medium parameters
- 2. Seismic wavelength
- 3. Oversampling in space

 $\Delta z = 3m$



How accurate is the reconstructed wavefield value one depth level <u>BELOW</u> the receivers?

 $P = \nabla \cdot u$ $S = \nabla \times u$ $\tilde{u}_x^p = \frac{1}{|k|} \left(k_x^2 \tilde{u}_x + k_x k_z \tilde{u}_z \right)$ $\tilde{u}_z^p = \frac{1}{|k|} \left(k_z^2 \tilde{u}_z + k_x k_z \tilde{u}_x \right)$ $\tilde{u}_x^s = \frac{1}{|k|} \left(k_z^2 \tilde{u}_x - k_x k_z \tilde{u}_z \right)$ $\tilde{u}_z^s = \frac{1}{|k|} \left(k_x^2 \tilde{u}_z - k_x k_z \tilde{u}_x \right)$



<u>Content</u>

1. Motivation

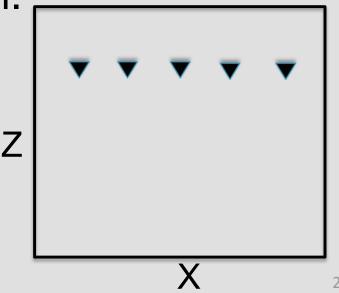
2. Theory

- General concept
- Inversion setup
- 3. Synthetic examples
 - Land
 - Ocean-bottom seismic
- 4. Conclusions
- 5. Road ahead

Observed Data: a shot gather consisting of:

- 1.Vertical geophone = Pup + Pdown + S + Surface + No-
- 2. Horizontal geophone = Pup + Pdown + S + Surface + No

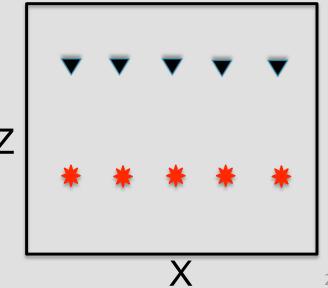
- Rough estimate of medium parameters at receiver's location.
- Constant throughout the inversion.



Model: a virtual source array consisting of:

- 1. Vertical displacement sources
- 2. Horizontal displacement sources

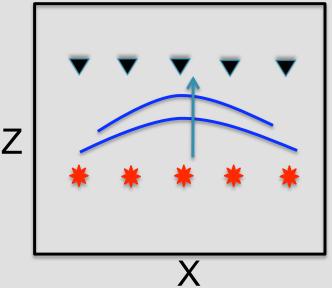
• Virtual sources located "somewhere" in medium.



Model: a virtual source array consisting of:

- 1. Vertical displacement sources
- 2. Horizontal displacement sources

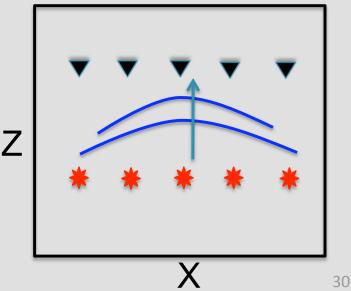
- Virtual sources located "somewhere" in medium.
- Virtual sources' functions are propagated; wavefield is recorded at receivers.



Model: a virtual source array consisting of:

- 1. Vertical displacement sources
- 2. Horizontal displacement sources

- Virtual sources located "somewhere" in medium.
- Virtual sources' functions are propagated; wavefield is recorded at receivers.
- Recorded vertical and horizontal displacements are
- "reconstructed data"



Inversion is complete when:

reconstructed displacements \approx observed displacements.

$$J = \frac{1}{2} \left\| F\vec{m} - \vec{g}_{obs} \right\|^2$$

$$ec{m}$$
 - Virtual source functions

- $F\,$ Propagation operator
- \vec{g} Geophone data

 $F\,$ - Elastic isotropic propagation operator

$$\nabla ((\lambda + \mu) \nabla \cdot \vec{u}) + \nabla \cdot (\mu \nabla \vec{u}) + \vec{f} = \rho \vec{\ddot{u}}$$

$$\begin{bmatrix} (\lambda + 2\mu)\partial_x^2 u_x + (\lambda + \mu)\partial_x \partial_z u_z + \mu \partial_z^2 u_x + f_x \\ (\lambda + 2\mu)\partial_z^2 u_z + (\lambda + \mu)\partial_x \partial_z u_x + \mu \partial_x^2 u_z + f_z \end{bmatrix} = \rho \begin{bmatrix} \partial_t^2 u_x \\ \partial_t^2 u_z \end{bmatrix}$$

 \vec{u} - Particle displacement

ho - Density $\mu =
ho V_S^2$ - Lame parameters $\lambda =
ho V_P^2 - 2\mu$

- ${\it F}$ Elastic isotropic propagation operator
 - ho_{P}, V_{P}, V_{s} Constant medium parameters. F is a linear operator.

After convergence:

1. Forward model using virtual source functions, to generate displacement fields: u(x,z)

$$d_x(x,z)$$

 $d_z(x,z)$

U

$$\begin{bmatrix} u_{x} \\ u_{z} \end{bmatrix} = F \begin{bmatrix} m_{x} \\ m_{z} \end{bmatrix}$$

 $u_{z}(x,z)$

After convergence:

1. Forward model using virtual source functions, to generate displacement fields: $u_x(x,z)$

$$\begin{bmatrix} u_{x} \\ u_{z} \end{bmatrix} = F \begin{bmatrix} m_{x} \\ m_{z} \end{bmatrix}$$

2. Apply divergence and curl operators on displacement fields at receiver depth:

$$P(x,z_{r}) = \frac{1}{2\Delta x} \left(u_{x}(x + \Delta x, z_{r}) - u_{x}(x - \Delta x, z_{r}) \right) + \frac{1}{2\Delta z} \left(u_{z}(x, z_{r} + \Delta z) - u_{z}(x, z_{r} - \Delta z) \right)$$
$$S(x,z_{r}) = \frac{1}{2\Delta x} \left(u_{z}(x + \Delta x, z_{r}) - u_{z}(x - \Delta x, z_{r}) \right) - \frac{1}{2\Delta z} \left(u_{x}(x, z_{r} + \Delta z) - u_{x}(x, z_{r} - \Delta z) \right)$$

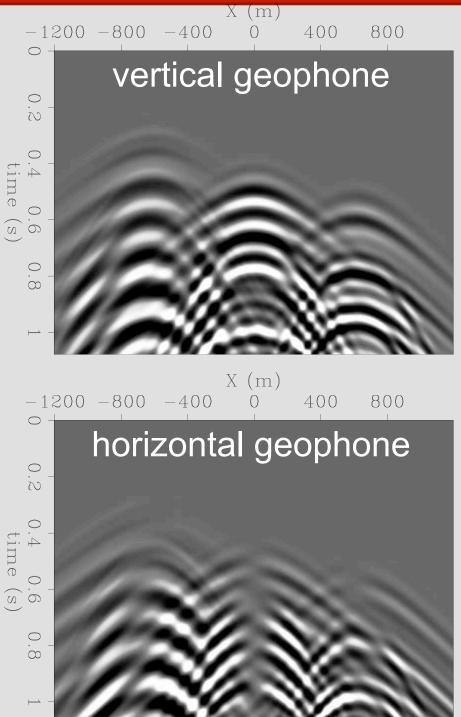
$$J = \frac{1}{2} \left\| F\vec{m} - \vec{g}_{obs} \right\|^2 \qquad \qquad \begin{bmatrix} u_x \\ u_z \end{bmatrix} = F \begin{bmatrix} m_x \\ m_z \end{bmatrix}$$

Objective function matches geophone data only.
Sufficient for P/S separation of land data.
Necessary but insufficient for P/S separation of OBS data.

<u>Content</u>

- 1. Motivation
- 2. Theory
 - General concept
 - Inversion setup
- 3. Synthetic examples
 - Land
 - Ocean-bottom seismic
- 4. Conclusions
- 5. Road ahead

Synthetic land data

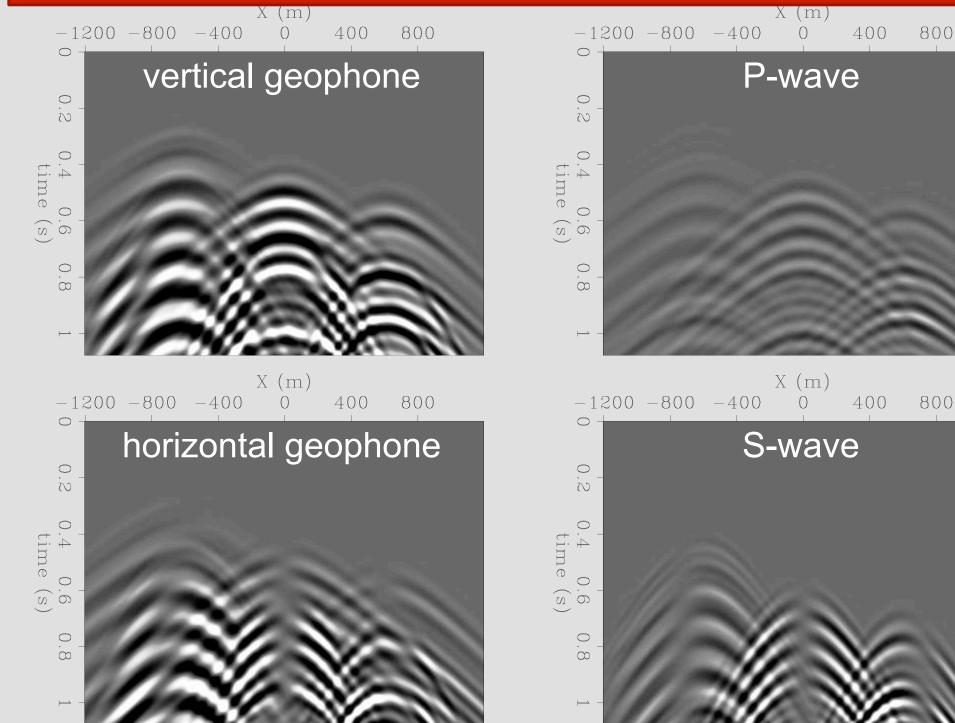


- •P and S waves.
- •Random reflectivity series.
- •Upgoing data only.
- •Sum of three shot gathers. with different V_p and V_s:

1.
$$v_{p} = 2.0 \frac{\text{km}}{\text{s}}, \quad v_{s} = 1.0 \frac{\text{km}}{\text{s}}, \quad \rho = 2.0 \frac{\text{gr}}{\text{cm}^{3}}$$

2. $v_{p} = 1.8 \frac{\text{km}}{\text{s}}, \quad v_{s} = 0.9 \frac{\text{km}}{\text{s}}, \quad \rho = 1.8 \frac{\text{gr}}{\text{cm}^{3}}$
3. $v_{p} = 1.6 \frac{\text{km}}{\text{s}}, \quad v_{s} = 0.8 \frac{\text{km}}{\text{s}}, \quad \rho = 1.6 \frac{\text{gr}}{\text{cm}^{3}}$

Synthetic land data



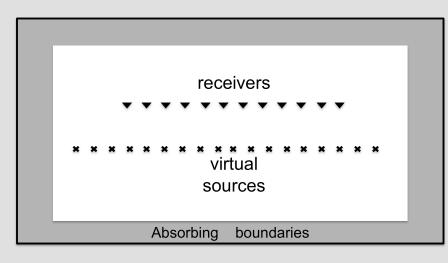
Synthetic land data

Synthetics generated with:

1.
$$v_p = 2.0 \frac{\text{km}}{\text{s}}, \quad v_s = 1.0 \frac{\text{km}}{\text{s}}, \quad \rho = 2.0 \frac{\text{gr}}{\text{cm}^3}$$

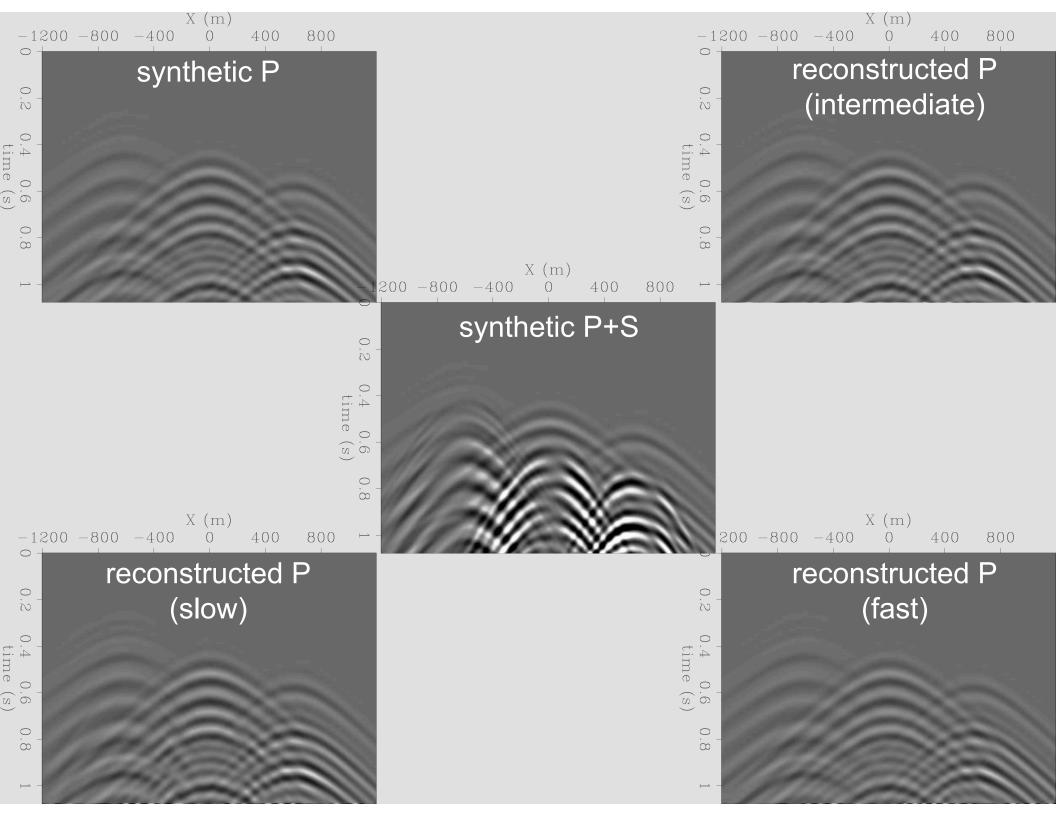
2. $v_p = 1.8 \frac{\text{km}}{\text{s}}, \quad v_s = 0.9 \frac{\text{km}}{\text{s}}, \quad \rho = 1.8 \frac{\text{gr}}{\text{cm}^3}$
3. $v_p = 1.6 \frac{\text{km}}{\text{s}}, \quad v_s = 0.8 \frac{\text{km}}{\text{s}}, \quad \rho = 1.6 \frac{\text{gr}}{\text{cm}^3}$

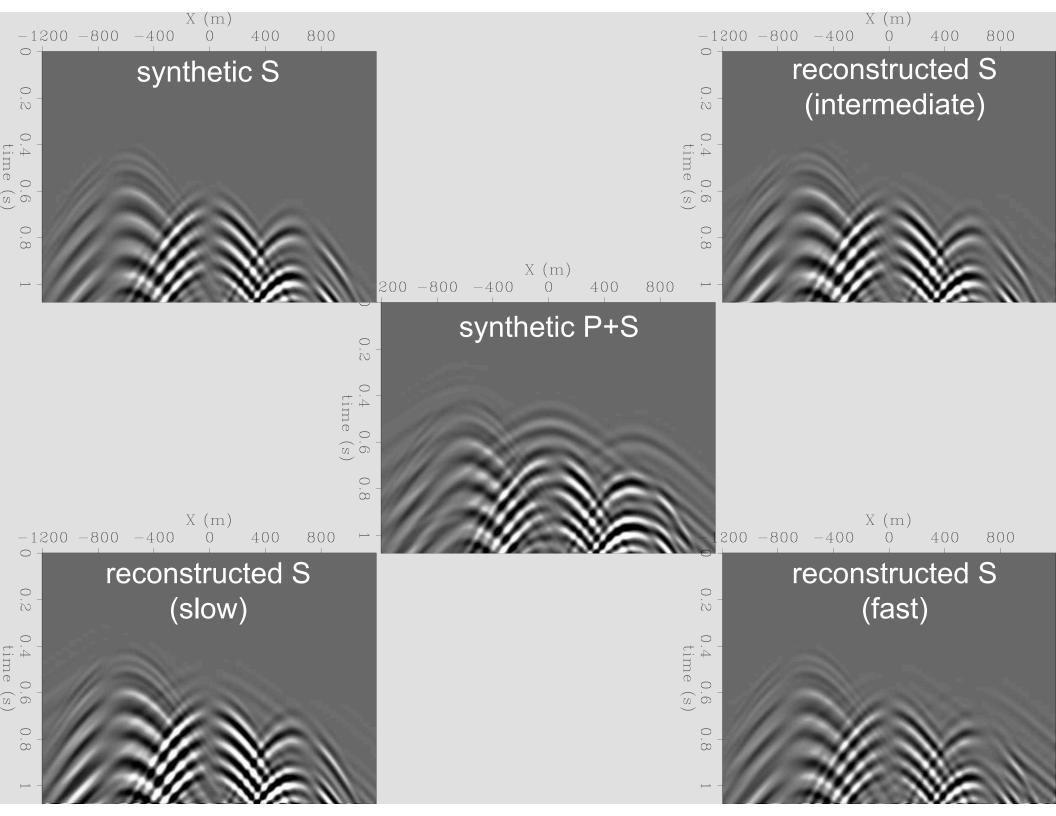
Three separate inversions with:



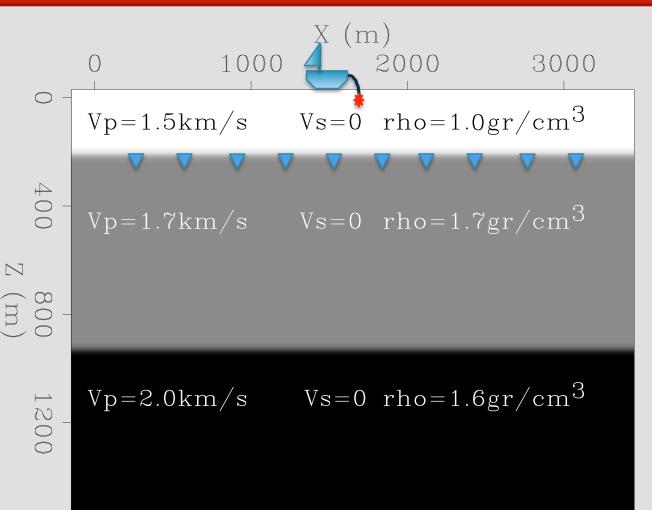
1. intermediate
$$v_{p} = 1.7 \frac{\text{km}}{\text{s}}, \quad v_{s} = 0.85 \frac{\text{km}}{\text{s}}, \quad \rho = 1.7 \frac{\text{gr}}{\text{cm}^{3}}$$

2. slow $v_{p} = 1.5 \frac{\text{km}}{\text{s}}, \quad v_{s} = 0.7 \frac{\text{km}}{\text{s}}, \quad \rho = 1.5 \frac{\text{gr}}{\text{cm}^{3}}$
3. fast $v_{p} = 2.0 \frac{\text{km}}{\text{s}}, \quad v_{s} = 1.0 \frac{\text{km}}{\text{s}}, \quad \rho = 2.0 \frac{\text{gr}}{\text{cm}^{3}}$





Synthetic OBS data

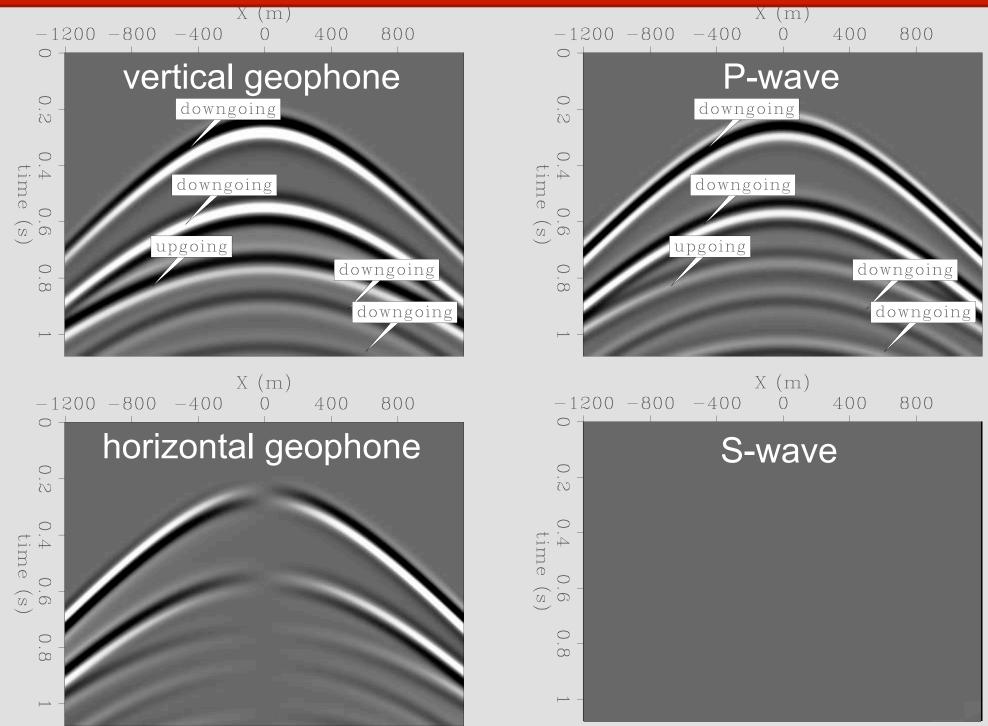


Inversion medium parameters:

$$v_{p} = 1.7 \frac{\text{km}}{\text{s}}$$
$$v_{s} = 0.85 \frac{\text{km}}{\text{s}}$$
$$\rho = 1.7 \frac{\text{gr}}{\text{cm}^{3}}$$

Direct arrival muted.

Synthetic OBS data



Synthetic OBS data X (m) X (m) -1200 - 800 - 400400 800 -1200 - 800 - 4000 400 800 0 \bigcirc synthetic P reconstructed P downgoing \bigcirc . N 0.4).4 0.6 time (s) downgoing pgoing downgoing 0.8 downgoing X (m) X (m) -1200 - 800 - 4000 400 800 -1200 - 800 - 4000 400 800 \bigcirc reconstructed S synthetic S \bigcirc .. N 0.4 time 0.6 (S)0. 8

 \bigcirc

0.2

0.4

0.6

0.8

 \bigcirc

 \bigcirc

. N

0.4

 \bigcirc

 \bigcirc

 ∞

 \mapsto

(S) (S)

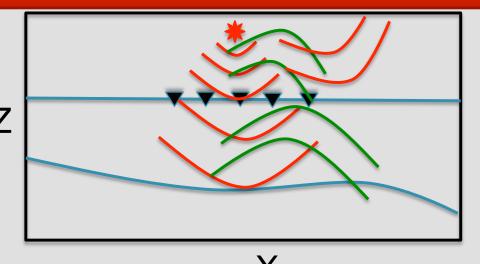
time

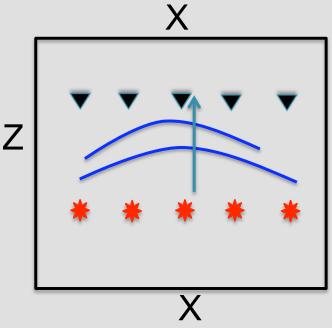
time (s)

Null-space issue

<u>Field data:</u> Upgoing and downgoing waves. Z

Reconstructed data: Upgoing waves only.



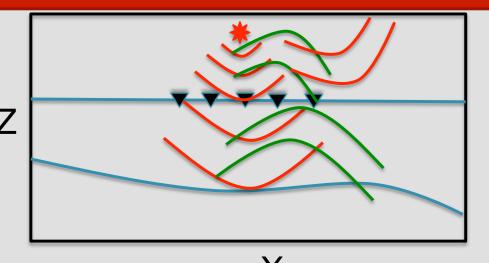


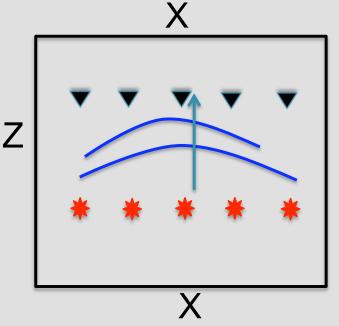
Null-space issue

<u>Field data:</u> Upgoing and downgoing waves. Z

Reconstructed data: Upgoing waves only.

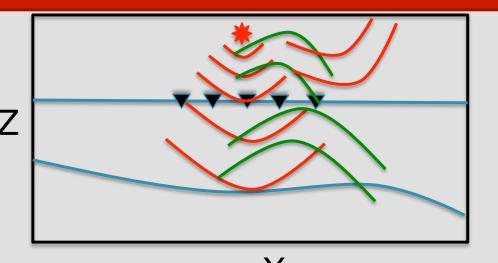
Vertical derivative is unconstrained. P and S are unconstrained.



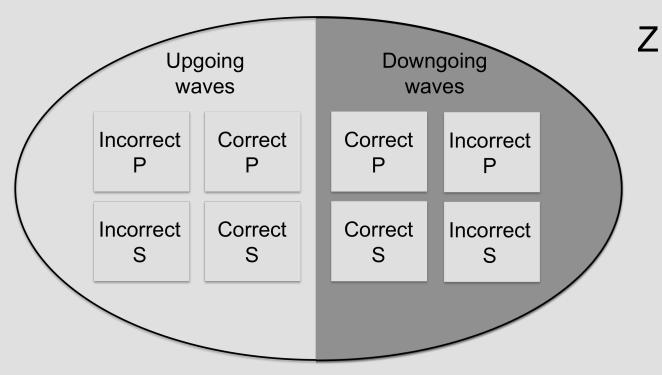


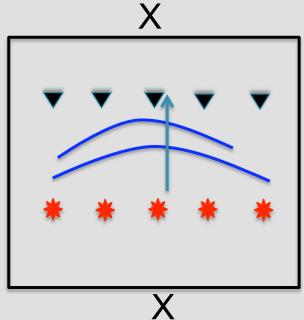
Null-space issue

<u>Field data:</u> Upgoing and downgoing waves. Z



Reconstructed data: Upgoing waves only.





Managing the solution-space and the null-space

Controls:

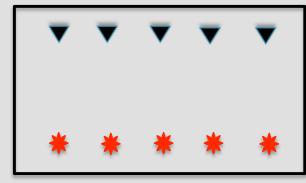
- 1. Virtual-source locations.
- 2. Additional constraints.

Question: Can the solution-space and the null-space be controlled in a useful way?

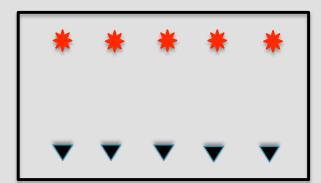
Example: prevent leakage of P-wave energy to S-wave energy.

Virtual-source locations

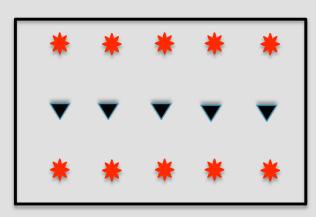
Control which wavenumbers can be reconstructed.



Upgoing only



Downgoing only

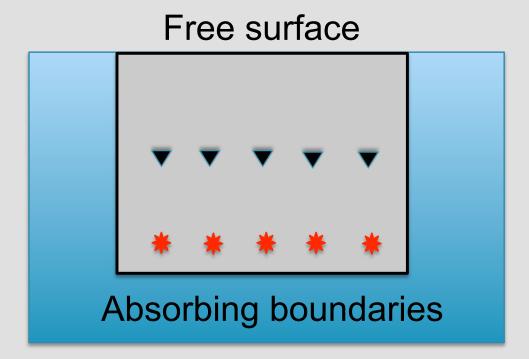


Increases null-space.

Upgoing and downgoing

Virtual-source locations

Use a free upper boundary to generate the downgoing waves.



Observed hydrophone *H*.

$$H \Leftrightarrow P = \frac{1}{2} (\lambda + \mu) \nabla \cdot \vec{u}$$
$$\vec{u} \Leftarrow \frac{1}{2} (\lambda + \mu) \nabla P$$

Observed hydrophone *H*.

$$H \Leftrightarrow P = \frac{1}{2} (\lambda + \mu) \nabla \cdot \vec{u} \qquad P = \frac{1}{2} (\lambda + \mu) \nabla \cdot \vec{u}^{P}$$
$$\vec{u} \leftarrow \frac{1}{2} (\lambda + \mu) \nabla P \qquad \vec{u}^{P} \leftarrow \frac{1}{2} (\lambda + \mu) \nabla P$$

Observed hydrophone *H*.

$$H \Leftrightarrow P = \frac{1}{2} (\lambda + \mu) \nabla \cdot \vec{u} \qquad P = \frac{1}{2} (\lambda + \mu) \nabla \cdot \vec{u}^{P}$$
$$\vec{u} \leftarrow \frac{1}{2} (\lambda + \mu) \nabla P \qquad \vec{u}^{P} \leftarrow \frac{1}{2} (\lambda + \mu) \nabla P$$

Modified objective function:

$$J = \varepsilon_1 \left\| F \vec{m} - \vec{g}_{obs} \right\|^2 + \varepsilon_2 \left\| A \vec{m} - H_{obs} \right\|^2$$

 $A = \nabla \cdot F$ $\left(\lambda + \mu\right) \Leftrightarrow \varepsilon_{2}$

Observed hydrophone *H*

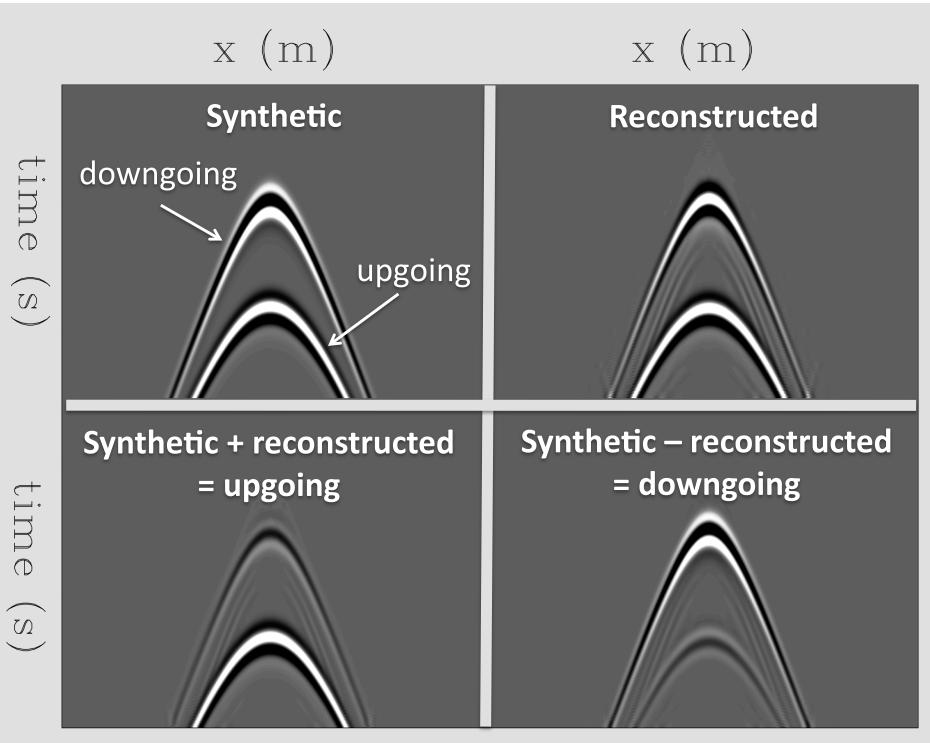
$$H \Leftrightarrow P = \frac{1}{2} (\lambda + \mu) \nabla \cdot \vec{u} \qquad P = \frac{1}{2} (\lambda + \mu) \nabla \cdot \vec{u}^{P}$$
$$\vec{u} \Leftarrow \frac{1}{2} (\lambda + \mu) \nabla P \qquad \vec{u}^{P} \Leftarrow \frac{1}{2} (\lambda + \mu) \nabla P$$

2

Modified objective function:

$$J = \varepsilon_1 \|F\vec{m} - \vec{g}_{obs}\|^2 + \varepsilon_2 \|A\vec{m} - H_{obs}\|$$
$$A = \nabla \cdot F$$
$$(\lambda + \mu) \Leftrightarrow \varepsilon_2$$
Cor

Constrains spatial derivative of P particle motion.



vertical geophone

<u>Content</u>

- 1. Motivation
- 2. Theory
 - General concept
 - Inversion setup
- 3. Synthetic examples
 - Land
 - Ocean-bottom seismic
- 4. Conclusions
- 5. Road ahead

P/S separation of ocean-bottom seismic data by inversion

- It is possible to generate equivalent seismic data using virtual-sources and incorrect medium parameters.
- P/S separation is reasonably good even for large medium parameter error.
- Reconstructable wavenumbers depend on virtual sources' locations.
- More constraints are required so that the spatial derivatives of the reconstructed field near the receivers are correctly estimated.

Objective: Solve the OBS problem

- 1. Utilise different virtual-source / receiver arrangements to enable reconstruction of all wavenumbers (increase null-space).
- 2. Combine geophone and hydrophone constraints (reduce null-space).

P/S separation of ocean-bottom seismic data by inversion

Thanks for listening